# A comparison between ES-MDA and restart EnKF for the purpose of the simultaneous identification of a contaminant source and hydraulic conductivity

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# Abstract

Understanding a contaminant source may help in a better management and risk assessment of a polluted aquifer. However, contaminant source information may not be available when a pollutant is detected in a drinking well. The restart ensemble Kalman filter (restart EnKF, also named r-EnKF) has been demonstrated in synthetic and laboratory experiments as an efficient solution for the identification of a contaminant source. Recently, the ensemble smoother with multiple data assimilation (ES-MDA) has been proposed as an alternative to the r-EnKF as a more efficient solution given that the r-EnKF needs to restart the simulation of the state equation from time zero after each data assimilation step. An analysis, in a synthetic aquifer, of the accuracy of the ES-MDA for the simultaneous identification of a contaminant source and the spatial distribution of hydraulic conductivity by assimilating both piezometric head and concentration observations is carried out using the r-EnKF as a benchmark. The conclusion is that the ES-MDA can outperform the r-EnKF, but the expected speed advantage, associated with the possibility of assimilating all data at once, does not exist. For the ES-MDA to reach the same level of accuracy as the r-EnKF, the

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number of multiple data assimilations must be large, and final computing time is similar for both approaches. However, the ES-MDA can do much better than the r-EnKF if the number of iterations increases even further, with the consequent increase of computational cost. *Keywords:* Contaminant source identification; Data assimilation; Ensemble smoother with multiple data assimilation; Restart ensemble Kalman filter

#### 1 1. Introduction

When a contaminant is released into the subsurface, it will jeopardize not only human health but also damage the local ecosphere, especially if the contaminant is hazardous. When contamination happens inadvertently or is purposely hidden, it may be difficult to trace it back from concentration observations taken downstream from the source. Yet, knowledge of the contaminant source is vital for groundwater contamination management, contamination control, contamination risk assessment and remediation.

<sup>8</sup> How to identify a contaminant source once contamination has been detected has attracted <sup>9</sup> much attention in the last decades. It is an intricate problem that has been addressed <sup>10</sup> using inverse modeling. According to their characteristics, the inverse modeling approaches <sup>11</sup> for contaminant source identification could be classified into three categories: optimization <sup>12</sup> approaches, probabilistic approaches, and deterministic approaches. The reader is referred to <sup>13</sup> the reviews by Sun et al. (2006a); Atmadja and Bagtzoglou (2001b); Michalak and Kitanidis <sup>14</sup> (2004); Bagtzoglou and Atmadja (2005) for further information.

In the optimization approaches, the objective is to minimize an objective function that measures the differences between simulated concentrations and measurement observations and that is written in terms of the parameters defining the contaminant source. Some of the approaches used are least-squares regression and linear programming (Gorelick et al., 1983), maximization of correlation coefficients (Sidauruk et al., 1998), constrained robust least squares (CRLS) (Sun et al., 2006a), CRLS estimation combined with a branch-and<sup>21</sup> bound global optimization (Sun et al., 2006b), evolutionary search algorithms (Mirghani
<sup>22</sup> et al., 2009), or hybrid simulation-optimization (Ayvaz, 2016).

In the probabilistic approaches, the objective is, generally speaking, to maximize some posterior probability of the source parameters given the observations. Some approaches used for this purpose are minimum relative entropy (Woodbury and Ulrych, 1996; Woodbury et al., 1998; Cupola et al., 2015), the geostatistical approach (Sun, 2007; Gzyl et al., 2014; Butera et al., 2013), Markov chain Monte Carlo (Wang and Jin, 2013), or Bayesian approaches (Zeng et al., 2012; Zhang et al., 2015; Zanini and Woodbury, 2016).

In the deterministic approaches, the main objective is to solve the advection-dispersion equation backward in time. Some of the approaches employ the marching-jury backward beam equation method (Atmadja and Bagtzoglou, 2001a; Bagtzoglou and Atmadja, 2003), Tikhonov regularization (Skaggs and Kabala, 1994; Neupauer et al., 2000), or a quasireversibility method together with minimum relative entropy (e.g., Skaggs and Kabala, 1995; Bagtzoglou and Atmadja, 2003; Neupauer et al., 2000).

In addition to the approaches mentioned above, recently, the use of the restart ensem-35 ble Kalman filter (r-EnKF) (a probabilistic approach), was proposed by Xu and Gómez-36 Hernández (2016) to identify a contaminant source by assimilating concentration observa-37 tions. The good results obtained by the r-EnKF in standard inversion problems (e.g., Hen-38 dricks Franssen and Kinzelbach, 2009; Xu et al., 2013; Xu and Gómez-Hernández, 2015b) 39 prompted its use for source identification, where it proved to achieve good results, too. Later, 40 Xu and Gómez-Hernández (2018) extended their work to jointly identify the source infor-41 mation and the underlying hydraulic conductivity field in a synthetic aquifer, and in a tank 42 experiment (Chen et al., 2018). Their works have proven the capability of the r-EnKF for 43 contaminant source identification. 44

The ensemble smoother (ES), also a probabilistic approach, first proposed by Van Leeuwen and Evensen (1996), is an alternative that could alleviate the computational burden of the EnKF, because it assimilates all data for all time steps at once. This avoids the restart of the simulation at every time step and makes the ES faster and easier to implement than the EnKF (Emerick and Reynolds, 2013a). However, the performance of the ES for the case of non-linear state equations is not good (e.g., Evensen and Van Leeuwen, 2000; Crestani et al., 2013), the main reason being the lack of multiple updatings inherent to the EnKF (the ES does only one update).

A detailed explanation of why the EnKF outperforms the ES in dealing with non-linear 53 problems can be found in the work by Evensen (2018). Here, a brief explanation is given. 54 The updating step in both the EnKF and the ES are written in terms of covariances, which 55 can only capture linear relationships. The EnKF recursively updates the parameters of 56 interest by assimilating observation information in time and after each step the updates get 57 closer to the reference solution. The ES makes a single update using all the data from all 58 time steps. That is, the EnKF is equivalent to making many linear approximations to the 59 state equation followed by incremental updates along the linear approximation, whereas the 60 ES is equivalent to a single linear approximation to the state equation and a single large 61 update along the linear approximation. Therefore, the EnKF is equivalent to a non-linear 62 optimization based on local linear approximations, whereas the ES is a linear minimization, 63 which may be very far from optimal if the state equation is highly nonlinear. Unless, iteration 64 is also introduced into the ES. This is what Emerick and Reynolds (2013a) propose with 65 their ensemble smoother with multiple data assimilation (ES-MDA). The basic idea is to 66 assimilate all data from all time steps several times, progressively updating the parameters 67 after each iteration. 68

Several successful applications of the ES-MDA are reported in the reservoir historymatching literature (e.g., Emerick et al., 2013; Emerick and Reynolds, 2013b; Le et al., 2015, 2016; Lee et al., 2013; Fokker et al., 2016). In these works, the reservoir state equations are nonlinear, and the ES-MDA results outperforms the EnKF for both synthetic and real <sup>73</sup> field problems. Recently, a few applications have been reported in the hydrogeology litera<sup>74</sup> ture (Li et al., 2018a,b) for the characterization of hydraulic conductivities by assimilating
<sup>75</sup> piezometric heads.

In this paper, the ES-MDA is used, for the first time, to the best of our knowledge, to jointly identify a heterogeneous hydraulic conductivity field and contaminant source information on a synthetic aquifer. As a benchmark, the accuracy of the ES-MDA will be compared with the r-EnKF. Note that the main aim of this work is to evaluate the capabilities of the ES-MDA and to benchmark it against the r-EnKF for the joint identification of conductivity field and contaminant source information.

The paper is organized as follows. First, we introduce the algorithmic description of the r-EnKF and the ES-MDA. Second, we test and compare the ES-MDA with the r-EnKF on a synthetic aquifer. And third, we discuss the results.

# 85 2. Methodology

# 86 2.1. Restart ensemble Kalman filter

The EnKF was developed based on the Kalman filter proposed by Kalman et al. (1960) 87 to better tackle nonlinear state-transfer equations. The main difference between the EnKF 88 and the Kalman filter is on how the covariance matrices are calculated. In the original filter, 89 the covariances were propagated in time using a linear state-transfer function (or a linear 90 approximation in case the function is non-linear), while in the EnKF, the covariances are 91 calculated from the states obtained after solving the state-transfer function on an ensemble 92 of realizations (e.g., Evensen, 2003, 2009; Chen and Zhang, 2006; Xu et al., 2013; Xu and 93 Gómez-Hernández, 2015a). Like the Kalman filter, the EnKF consists of two steps: forecast 94 and analysis. The first one is to forecast the state variables from the state variables and 95 the best estimate of the model parameters from the last time step. And the second one 96 is to update the state variables and model parameters at the current time step based on 97

the deviations between forecasted and observed state variable values at selected observation points. However, as already discussed in Xu and Gómez-Hernández (2016), it is impossible to take into account the updated parameters in the forecast step when these parameters define the spatiotemporal position of a contaminant source, unless the forecast is restarted from time zero. This approach modifies the standard Kalman filter equations since there is no need to update the variable values at the analysis step: their values will be recomputed with the new estimates of the model parameters from times zero.

For any given realization of the ensemble, let  $V_t^f$  denote the forecasted state variables at time t, and  $S_t^a$  the best model parameter estimates after the analysis step at the same time. The forecast equation is

$$V_t^f = \psi(V_0, S_{t-1}^a).$$
(1)

where  $\psi$  represents the state-transfer function, and  $V_0$  represents the state variables at time zero. The update step modifies the parameter values from the previous time step  $(S_{t-1}^a)$  as a function of the discrepancy between forecasted and observed state variables at observation locations

$$S_t^a = S_{t-1}^a + G_t^f(V_{o,t} + e_t - V_{o,t}^f)$$
(2)

112 with

$$G_t^f = D_{SV,t}^f (D_{VV,t}^f + R_t)^{-1}, (3)$$

where  $V_{o,t} + e_t$  is the vector of observed concentrations and piezometric heads (composed of the sum of the true head or concentration  $V_{o,t}$  plus an observation error  $e_t$  of zero mean and covariance  $R_t$ ),  $G_t^f$  is the Kalman gain,  $D_{SV,t}^f$  is the cross-covariance between parameters and forecasted state variables at observation locations, and  $D_{VV,t}^f$  is the auto-covariance between the forecasted state variables at the observation locations.

Consider that there are  $N_r$  realizations in the ensemble and each realization has been discretized into  $N_e$  elements. The state variable vector V contains piezometric heads H and  $_{120}$  concentrations C at all aquifer model cells

$$V = \begin{bmatrix} H \\ C \end{bmatrix}.$$
 (4)

# This vector contains $N_r$ realizations of $2N_e$ variables.

The model parameter vector S contains hydraulic log-conductivity  $\ln K$  in all aquifer model cells and the contaminant source parameters, which are source location, X for the x-coordinate, and Y for the y-coordinate, initial release time T, release duration  $\Delta T$ , and mass-loading rate M

$$S = \begin{bmatrix} \ln K \\ X \\ Y \\ T \\ \Delta T \\ M \end{bmatrix}.$$
(5)

This vector contains  $N_r$  realizations of  $(N_e + 5)$  variables.

Then, if we define  $d_t = V_{o,t} + e_t - V_{o,t}^f$  and  $P_{VV,t}^f = (D_{VV,t}^f + R_t)^{-1}$ , and the covariances are split into the auto- and cross- covariances of each parameter, the updating equation (2), applicable to each realization independently, can be written as

$$S_{t}^{a} = \begin{pmatrix} \ln K \\ X \\ Y \\ Y \\ T \\ \Delta T \\ M \end{pmatrix} + \begin{pmatrix} D_{(\ln K)C,t}^{f} & D_{(\ln K)H,t}^{f} \\ D_{XC,t}^{f} & D_{XH,t}^{f} \\ D_{YC,t}^{f} & D_{YH,t}^{f} \\ D_{TC,t}^{f} & D_{TH,t}^{f} \\ D_{TC,t}^{f} & D_{TH,t}^{f} \\ D_{(\Delta T)C,t}^{f} & D_{(\Delta T)H,t}^{f} \\ D_{MC,t}^{f} & D_{MH,t}^{f} \end{pmatrix} \begin{pmatrix} P_{CC,t}^{f} & P_{CC,t}^{f} \\ P_{HC,t}^{f} & P_{HC,t}^{f} \end{pmatrix} \begin{pmatrix} d_{C,t} \\ d_{H,t} \end{pmatrix}$$
(6)

<sup>130</sup> 2.2. Ensemble smoother with multiple data assimilation

The ES is, conceptually, the same as the r-EnKF but limited to one forecast step (for all the time steps for which observations are available) and a single update step (based on the discrepancies between observations and predictions at all time steps).

The equations that describe the ES are almost the same as those for the r-EnKF above, with some differences. The forecast step is given by

$$V^{f} = \psi(V_0, S_0).$$
(7)

where now  $V^f$  contains the state forecasted at all time steps —computed from the initial state  $V_0$  and the initial ensemble of parameters  $S_0$ . And the update step is given by

$$S^{a} = S_{0} + G^{f}(V_{o} + e - V_{o}^{f}),$$
(8)

138 with

$$G^{f} = D^{f}_{SV} (D^{f}_{VV} + R)^{-1}, (9)$$

where  $V_o + e$  are all of the observations at observation locations, e are the observation 139 errors, and  $V_o^f$  are the forecasts at observation locations. The covariances appearing in Eq. 140 (9),  $D_{SV}^{f}$  and  $D_{VV}^{f}$  are computed for all time steps; these covariance matrices include the 141 cross-covariances between time steps, an aspect not accounted for in the r-EnKF that might 142 render the ES superior to the r-EnKF. From a computational point of view, if there are  $N_o$ 143 observations locations sampled  $N_t$  times, the sizes of the matrices involved in the r-EnKF are 144 proportional to  $N_o$ , whereas in the ES they are proportional to the product  $N_o \cdot N_t$ . Hence, 145 the sizes of the cross-covariances in the r-EnKF are  $(N_e + 5) \times 2N_o$  for  $D_{SV,t}^f$ , and  $2N_o \times 2N_o$ 146 for  $D_{VV,t}^f$  and  $R_t$ ; whereas the size of the cross-covariance for the ES are  $(N_e + 5) \times (2N_o \cdot N_t)$ 147 for  $D_{SV}^f$  and  $(2N_o \cdot N_t) \times (2N_o \cdot N_t)$  for  $D_{VV}^f$  and R. 148

As we stated before, the performance of the ES is not good when dealing with non-149 linear problems. The solution provided by Emerick and Reynolds (2013a) to improve the 150 performance of the ES for non-linear state-transfer equations is to iterate, what is called 151 multiple data assimilation (because the same data is assimilated multiple times) on the basis 152 that each iteration of the ES is similar to a Gauss-Newton iteration (Reynolds et al., 2006; 153 Gu and Oliver, 2007). Basically, Eq. (7) and Eq. (10) are iteratively applied using the 154 latest updated parameters as the initial parameters for the next iteration. However, since 155 all data are assimilated multiple times, there is a need to inflate the observation error for 156 each assimilation step. For this purpose, a non-increasing sequence of error variance inflation 157 coefficients  $\{a_i, i = 1, ..., N_a\}$  is used in the updating equations, with  $N_a$  being the number 158 of assimilation iterations, and satisfying that  $\sum_{i=1}^{N_a} \frac{1}{a_i} = 1$ . 159

<sup>160</sup> The ES-MDA equations display the following differences. The forecast step is given by

$$V_i^f = \psi(V_0, S_i^a).$$
(10)

where i is the iteration counter, and for each iteration the forecast uses the last updated parameters from the previous iteration. And the update equation is given by

$$S_i^a = S_{i-1}^a + G_i^f(V_{o,i} + \sqrt{a_i}e_i - V_{o,i}^f)$$
(11)

163 with

$$G_i^f = D_{SV,i}^f (D_{VV,i}^f + a_i R_i)^{-1}, (12)$$

In Eq. (11) and Eq. (12), we can see how the observation variance is amplified by a factor  $a_i$  and the observation error is amplified by  $\sqrt{a_i}$ .

If we define  $d_i = V_{o,i} + \sqrt{a_i}e_i - V_{o,i}^f$  and  $P_{VV,i}^f = (D_{VV,i}^f + a_iR_i)^{-1}$ , and the covariances are split into the auto- and cross- covariances of each parameter, the updating equation Eq.(8) <sup>168</sup> can be written as

$$S_{i}^{a} = \begin{pmatrix} \ln K \\ X \\ Y \\ T \\ \Delta T \\ M \end{pmatrix} + \begin{pmatrix} D_{(\ln K)C,i}^{f} & D_{(\ln K)H,i}^{f} \\ D_{XC,i}^{f} & D_{XH,i}^{f} \\ D_{YC,i}^{f} & D_{YH,i}^{f} \\ D_{TC,i}^{f} & D_{TH,i}^{f} \\ D_{(\Delta T)C,i}^{f} & D_{(\Delta T)H,i}^{f} \\ D_{MC,i}^{f} & D_{MH,i}^{f} \end{pmatrix} \begin{pmatrix} P_{CC,i}^{f} & P_{CC,i}^{f} \\ P_{HC,i}^{f} & P_{HC,i}^{f} \end{pmatrix} \begin{pmatrix} d_{C,i} \\ d_{H,i} \end{pmatrix}$$
(13)

Please notice that, when  $N_r < 2N_o$  in r-EnKF, or  $N_r < 2N_o \times N_t$  in ES-MDA, the low rank of the matrices prevent their inversion; then, the subspace inversion introduced by Evensen (2004) is used to solve for  $P_{VV,t}^f$  or  $P_{VV,i}^f$ . The detailed explanation can be found in the works by Evensen (2004); Emerick and Reynolds (2013a).

# 173 3. Application

A synthetic confined aquifer is designed and constructed on a 1000 [L] by 1000 [L] by 50 174 [L] prism discretized into 50 by 50 by 1 cells, where each cell is 20 [L] by 20 [L] by 50 [L]. 175 (Please note that no specific units are used throughout, only their dimensional analysis is 176 given. Any set of consistent units will yield the same results.) The reference log-conductivity 177 field is drawn from a multivariate Gaussian random function defined by the parameters in 178 Table 1 using the GCOSIM3D software —a sequential Gaussian simulation program (Gómez-179 Hernández and Journel, 1993). The resulting reference log-conductivity field is shown in 180 Figure 1. 181

Table 1: Parameters of the random functions used to generate the  $\ln K$  realizations. Spherical variogram with anisotropic spatial correlation defined by  $\lambda_{max}$  and  $\lambda_{min}$ , which are the ranges in the maximum and minimum directions of continuity. The angle corresponds to the maximum continuity direction and it is measured clockwise from the North direction

	Mean	Std. dev.	Variogram	$\lambda_{max}$	$\lambda_{min}$	Angle
$\ln K$	-1	1	Spherical	300	200	135

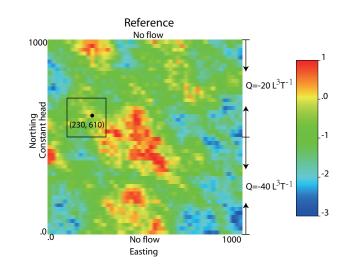


Figure 1: Reference  $\ln K$  and boundary conditions. The source location is marked with a dark dot. The inner square indicates the suspect contaminant source.

The model boundaries, as indicated in Figure 1, are set as follows: north and south boundaries are impermeable; west boundary is a prescribed head condition with a constant value of 50 [L]; east boundary is a prescribed flow boundary divided into two equal-length segments: the north segment with a total prescribed flow extraction rate of 20  $[L^{3}T^{-1}]$  and the south segment with a total extraction prescribed flow rate of 40  $[L^{3}T^{-1}]$ . Figure 2 shows the location of the 25 observation wells (red triangles) and the two verification wells (blue diamonds).

The initial concentration is zero  $[ML^{-3}]$  and the initial head for the whole domain is 58 [L], except at the west constant boundary. Other groundwater flow and contaminant transport

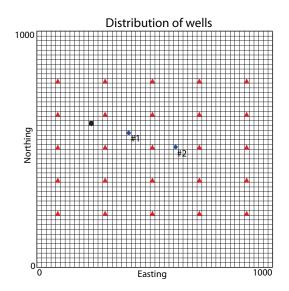


Figure 2: Location of wells. Red triangles mark observation wells; blue diamonds mark verification wells. The black circle is the contaminant source location.

parameters are assumed known and set as homogeneous: porosity of 0.3 [-], longitudinal
dispersivity of 2 [L], transverse to longitudinal dispersivity ratio of 0.1.

We assume the contaminants are inert. Only advection and dispersion are considered as transport mechanisms. Both groundwater flow and contaminant transport are under transient conditions. The groundwater flow simulator MODFLOW (McDonald and Harbaugh, 1988) and the transport simulator MT3DMS (e.g., Zheng, 2010; Ma et al., 2012) are used as forward models to solve the groundwater flow and contaminant transport problems, respectively.

The total simulation time is 10000 [T] and is discretized into 100 time steps with increasing size following a geometric series with ratio 1.01 (The first time step is 58.66 [T]). The observations of both piezometric head and concentration from the first 60 time steps (around 4790 [T]) are assimilated for the purpose of parameter identification, so the total number of observations is  $2 \times 25 \times 60$ .

The contaminant is released at location (X, Y) = (230, 610) [L] with a mass-loading rate of 1000 [MT<sup>-1</sup>], starting at time 613 [T] (around the 10th time step) and ending at time <sup>206</sup> 2867 [T] (around the 40th time step), with a release duration of 2254 [T].

Figure 3 shows three snapshots of piezometric head and solute concentration taken on the reference aquifer at the 10th simulation time step (beginning of contaminant injection), 40th time step (end of contaminant injection), and at 60th time step (end of assimilation period). This figure also shows the location where both piezometric heads and concentrations are sampled for the purpose of their assimilation in the different scenarios described next.

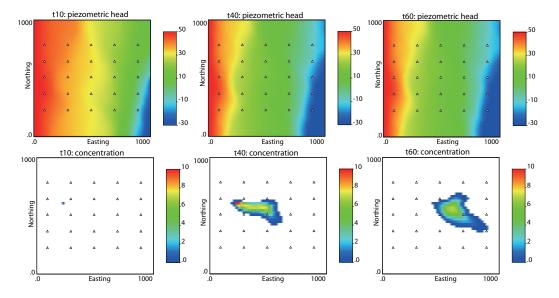


Figure 3: Reference. Piezometric head (top row) and contaminant plume (bottom row) at the 10th (beginning of solute injection), 40th (end of solute injection), and 60th (end of assimilation) time steps in the reference aquifer. White triangles mark the observation wells.

Seven scenarios will be evaluated. The first one, used as a benchmark to evaluate the 212 efficiency of the ES-MDA, is the r-EnKF, which has already proven its ability for the iden-213 tification of contaminant source parameters and hydraulic conductivity characterization; it 214 will be referred to as S0. The second one is the ES in its original implementation, that is, 215 without any iteration. Then, to evaluate the effect of the number of iterations, the ES-MDA 216 is run for five different scenarios, the difference between them is the number of iterations 217 (or data assimilations) performed; they will be labeled S2 to S6 with 2, 4, 6, 8 and 10 it-218 erations, respectively. Notice that the observation error inflation coefficients  $a_i$  will, in all 219

cases, be equal to the number of iterations, following the recommendations by Emerick and Reynolds (2013a), who show that the use of decreasing inflation coefficients leads to only small improvements with respect to using the inflation coefficients equal to the number of iterations.

As we stated before, the total number of cells of the log-conductivity field is  $50 \times 50$ , and 224 the number of source parameters is 5, so the total number of parameters to be identified is 225 2505. An ensemble of 400 initial log-conductivity realizations is generated using the same 226 random function model and parameters as for the reference log-conductivity field. The 227 number of ensemble members was chosen after a previous analysis with ensemble sizes of 228 200, 400 and 800 members. The difference in results between the ensemble sizes of 400 and 220 800 were not large enough to grant the use of the largest ensemble. Notice that there are 230 no conditioning log-conductivity data, thus the ensemble mean and ensemble variance of 231 the initial log-conductivity realizations are homogeneous and equal to their marginal values. 232 As already discussed by Xu et al. (2013) the use of the same random function parameters 233 for the generation of the initial realizations as for the generation of the reference case is 234 only a marginal advantage given that there are no conditioning conductivities. Indeed, 235 Xu et al. (2013) demonstrate the effectiveness of the r-EnKF using a totally uninformative 236 prior random function for the generation of the initial ensemble, with similar results as 237 when the "true" random function is used. In addition, an ensemble of 400 5-tuplets for the 238 source parameters is generated, each 5-tuplet contains five values drawn independently from 239 the following uniform distributions: initial release time  $T \in \mathcal{U}[550, 750]$ , release duration 240  $\Delta T \in \mathcal{U}[2100, 2300]$ , mass-loading rate  $M \in \mathcal{U}[900, 1100]$ , and source location  $(X, Y) \in \mathcal{U}[2100, 2300]$ 241  $(\mathcal{U}[100, 300] \times \mathcal{U}[500, 700]).$ 242

### 243 4. Results

Before starting the analysis of the results, Table 2 shows the CPU consumption for all 244 scenarios. Recall that in the r-EnKF (S0) there are 60 forecasting steps starting from time 245 0, and 60 assimilation steps to update the parameters 60 times based on the observations 246 at 25 wells; whereas, in the ES-MDA the number of model runs for the whole simulation 247 period is equal to the number of assimilation steps, but at each assimilation step, there are 248 1500 observations (25 observation locations times 60 time steps). For the current model and 249 setup, the ES-MDA is cheaper to run than the r-EnKF up until data are assimilated four 250 times. When ten iterations are performed, the ES-MDA costs two and half times that of the 251 r-EnKF. 252

Table 2: Definition of scenarios and CPU time consumption. The number in parenthesis refers to the number of data assimilation steps used in the ES-MDA. (ES would be equivalent to ES-MDA(1))

Method	Scenario	CPU in s	CPU in % of S0
r-EnKF	S0	16366	100%
ES	S1	4981	30%
ES-MDA(2)	S2	9526	58%
ES-MDA(4)	S3	17937	110%
ES-MDA(6)	S4	27432	149%
ES-MDA(8)	S5	34936	210%
ES-MDA(10)	S6	42422	259%

The r-EnKF, the ES and the ES-MDA will be used to assimilate the piezometric head and concentration data at the 25 observation locations. This assimilation will result in an ensemble of updated parameters (for the spatial distribution of hydraulic conductivity and for the parameters defining the contaminant source) that are used to produce an ensemble of piezometric heads and concentrations past the assimilation period (60th time step) for 40
time steps more. The performance of the different scenarios will be evaluated by comparing
the different final ensembles to their corresponding counterparts in the reference aquifer.

Figure 4 shows the ensemble mean and the ensemble variance of the updated log-conductivities 260 for scenarios S0 to S3 and S6. (The corresponding maps for S4 and S5 for this and following 261 figures are shown in the appendix.) The ensemble mean shows how the main patterns of 262 variability of the reference are captured by the updated ensemble, and the ensemble variance 263 shows the local variability of the updated log-conductivities. From a purely qualitative point 264 of view it is clear that the r-EnKF does a good job in capturing the reference patterns with 265 a small local uncertainty where the ensemble variance is close to zero, that the ES is able to 266 extract patterns which are, overall, similar to the reference but still far from them and with 267 a substantial local uncertainty, and that the ES-MDA gets better the more times data are 268 assimilated, with scenario S6 —for which data are assimilated 10 times —giving the best 269 results. 270

The above analysis can be quantified by computing the average absolute bias (AAB) and the ensemble spread (ESp). The AAB is used to measure the average absolute deviation between the updated values and the reference ones. The ESp measures the precision of the ensemble of updated realizations by calculating the root square of the ensemble variance. Their expressions are the following

$$AAB = \frac{1}{N_e} \sum_{i=1}^{N_e} \frac{1}{N_r} \sum_{j=1}^{N_r} |\ln K_{i,j} - \ln K_{i,ref}|, \qquad (14)$$

276

$$\mathrm{ESp} = \sqrt{\frac{1}{N_e} \sum_{i=1}^{N_e} \sigma_i^2},\tag{15}$$

where  $N_e$  is the number of model elements,  $N_r$  is the number of realizations,  $\ln K_{i,ref}$  is the reference log-conductivity value at node i,  $\ln K_{i,j}$  is the log-conductivity at node i for

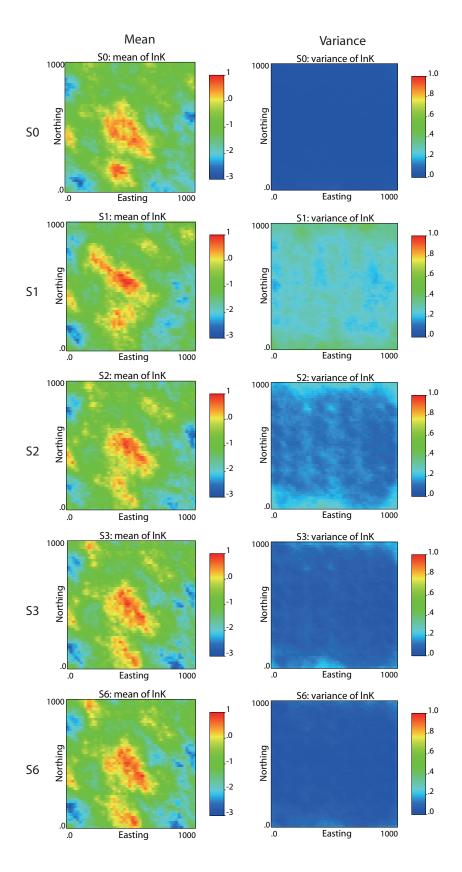


Figure 4: Scenarios S0-S3 and S6. Ensemble mean (left column) and ensemble variance (right column) of updated log-conductivity realizations. 17

realization j and  $\sigma_i$  is the log-conductivity ensemble variance at node i.

Figure 5 shows the AAB and ESp of lnK and of the parameters defining the contaminant 280 source for all scenarios, computed before any data assimilation and after data have been 281 assimilated over the first 60 time steps. The values, as expected, are the highest for the 282 initial ensembles. They are drastically reduced for the r-EnKF except for  $\Delta T$  and M. The 283 smoother provides increasingly smaller values as the number of assimilation steps increases, 284 with the best values for S6 after ten iterations. Specifically, the AAB and ESp of the updated 285 lnK, and Y for scenarios S3-S6 is close to that of scenario S0, and the AAB and ESp of the 286 updated T for scenario S6 is close to that of scenario S0; while, the AAB and ESp of the 287 updated X,  $\Delta T$  and M of scenarios S3-S6 is smaller than that of scenario S0. From these 288 results we could conclude that, after four assimilation steps, the ES-MDA starts to perform 289 better than the r-EnKF. 290

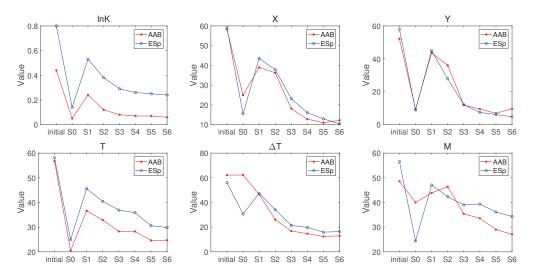


Figure 5: Scenarios S0-S6. Average absolute bias (AAB) and ensemble spread (ESp) of log-conductivity (lnK), source location (X and Y), initial release time (T), release duration  $(\Delta T)$ , and mass-loading rate (M) computed on the initial parameters and on the updated parameters for the different scenarios after 60 time steps.

Figure 6 shows the piezometric head distribution at the 60th time step computed with the final updated parameters for scenarios S0 to S3 and S6. The maps show, in the left column,

the piezometric head distributions for an individual ensemble member (realization #300), in 293 the center column, the ensemble mean obtained as the local mean of the piezometric head 294 at each node through the 400 realizations, and in the right column the ensemble variance. 295 Please, notice that the middle column with the ensemble mean piezometric heads is not the 296 solution of the state equations in the ensemble log-conductivity average of Fig. 4. An analysis 297 of these maps shows the robustness of the r-EnKF (S0) that produces an ensemble mean 298 map quite close to the reference one (upper right corner in Fig. 3) and with little variability 299 everywhere. The smoother performs well when comparing the average ensemble with the 300 reference map, but the uncertainties associated are quite large, especially in scenarios S1 and 301 S2: there is a need to assimilate the data at least four times (S3) to get a variance reduction 302 that approximates that of the r-EnKF. 303

Figure 7 shows the concentration plume computed with the parameters updated using 304 observations at 60 time steps. In the left column, the plume in realization #300, in the 305 center, the ensemble mean of the 400 plumes computed in the 400 realizations with updated 306 parameters, and in the right column the local concentration variance computed at each node 307 through the ensemble of realizations. Please, notice that, as with piezometric heads, the 308 middle column with the ensemble mean concentrations is not the solution of the state equa-309 tions in the ensemble log-conductivity average of Fig. 4. An analysis of these maps reaches 310 the same conclusions as for the piezometric heads, the r-EnKF is quite robust producing an 311 ensemble mean plume quite close to the reference (lower right corner in Fig. 3) and with 312 lower variability. The smoother performs well only when the number of iterations is large 313 (S3 and S6); for the cases of one, and two iterations (S1 and S2, respectively), the ensemble 314 mean plume is quite spread, the local variance is large, and the plume in the single selected 315 realization shown in the left column of the figure can be quite far from the reference one. 316

Figure 8 shows the time evolution of piezometric heads and solute concentrations at the two verification wells (#1 and #2) computed using the initial ensembles of contaminant

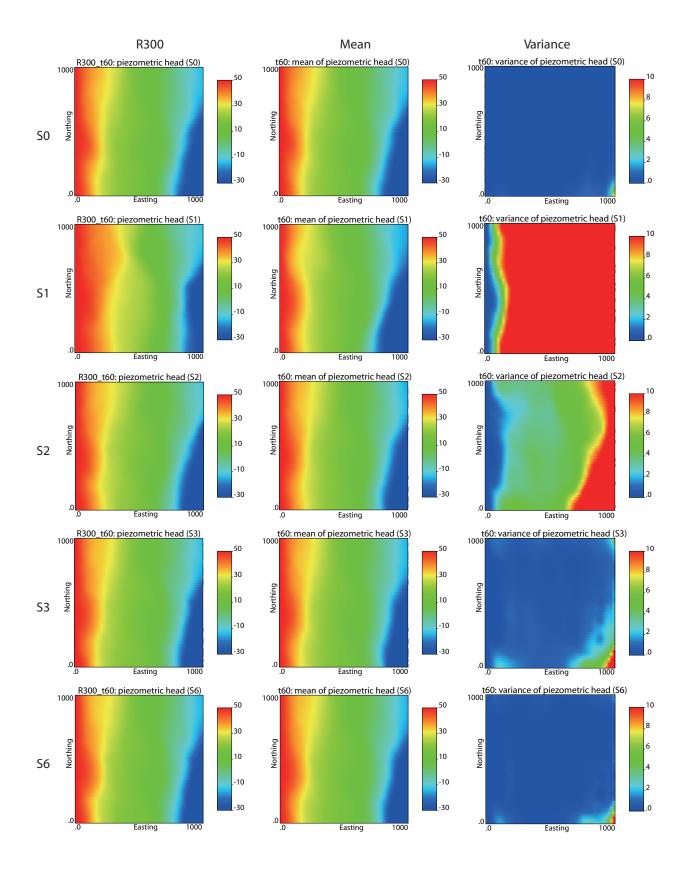


Figure 6: Scenarios S0-S3 and S6. Piezometric heads as computed with the updated parameters at the end of the 60th time step. From left to right, heads in read ation #300; ensemble mean, and ensemble variance.

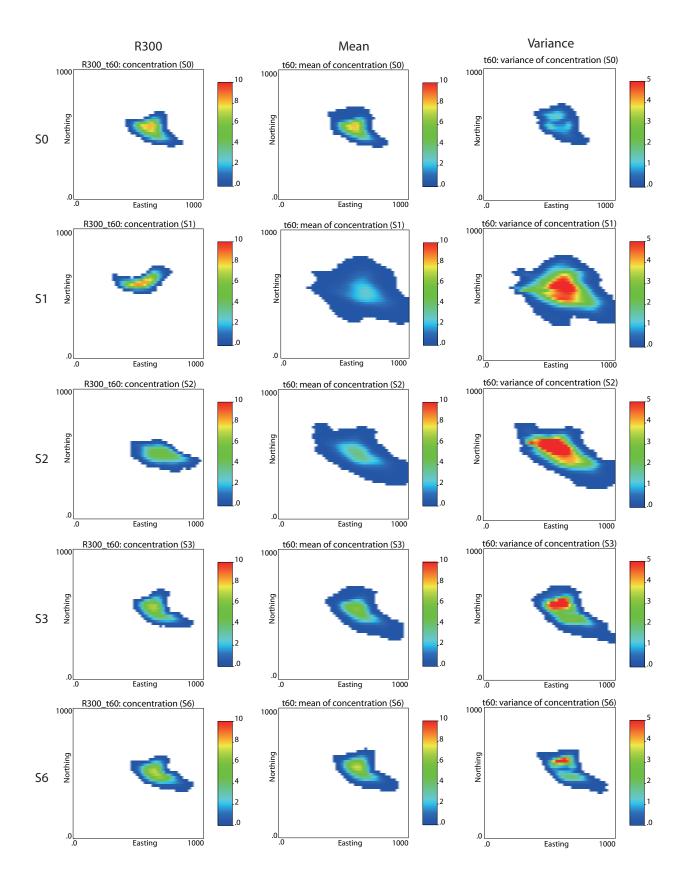


Figure 7: Scenarios S0-S3 and S6. Contaminant pl $\mathfrak{P}$  as computed with the updated parameters at the end of the 60th time step. From left to right, Contaminant plume in realization #300; ensemble mean of all contaminant plumes, and ensemble variance of all contaminant plumes.

source parameters and log-conductivities. The spread of predicted values is quite large 319 since no observation has been assimilated yet. Figure 9 and 10 show the time evolution of 320 piezometric heads and solute concentrations computed with the updated source parameters 321 and log-conductivity fields after the assimilation of the observations during the first 60 time 322 steps, respectively. The spread of the curves after the assimilation is considerably reduced, 323 especially for scenarios S0, S3 and S6. Although these two wells were not used during the 324 assimilation, the reproduction of piezometric heads, even after the assimilation period ends 325 is very good both for the r-EnKF (S0) and for the ES-MDA with four and ten iterations (S3 326 and S6), with the former performing slightly better than the latter. 327

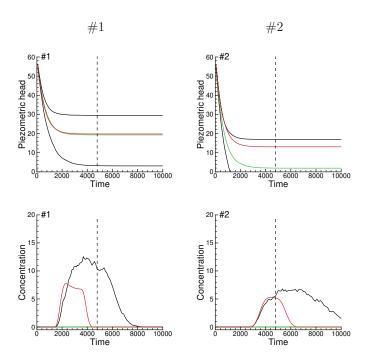


Figure 8: Time evolution of piezometric heads (top row) and solute concentrations (bottom row) at the two verification wells #1, and #2 computed on the initial ensemble of source information parameters and  $\ln K$ . The red line corresponds to the the reference field. The black lines correspond to the 5 and 95 percentiles of all realizations, and the green line corresponds to the median. The vertical dashed lines mark the end of the assimilation period.

<sup>328</sup> Up to here, regarding the characterization of the log-conductivity field and the repro-<sup>329</sup> duction of the state variables, the r-EnKF seems to outperform the ES-MDA with four

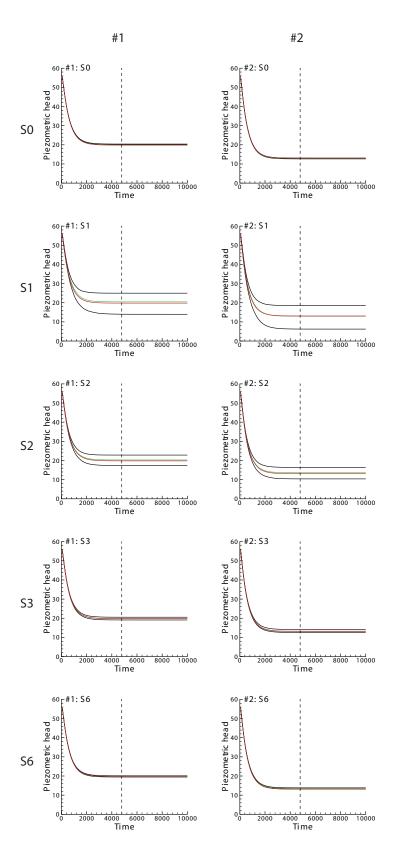


Figure 9: Scenarios S0-S3 and S6. Time evolution of the piezometric heads at the two verification wells #1, and #2 computed with the updated ensemble of source information parameters and  $\ln K$  after the assimilation of the observations of the first 60 time steps. The red line is the evolution of the piezometric head in the reference. The black lines correspond to the 5 and 95 percentiles of all realizations, and the green line corresponds to the median. The vertical dashed lines mark the end of the assimilation period.

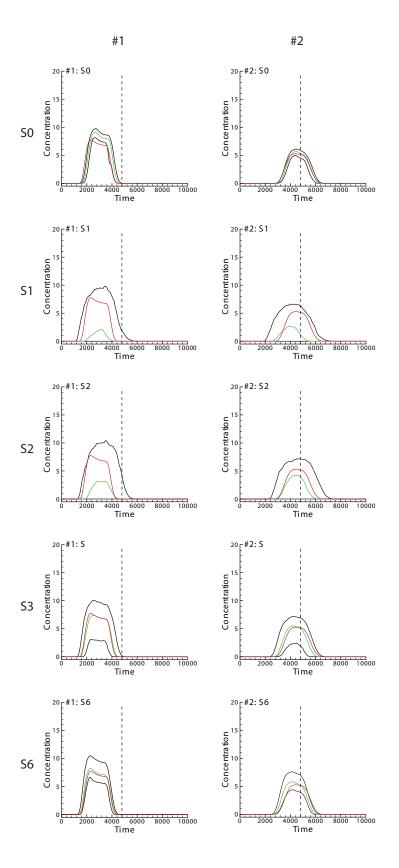


Figure 10: Scenarios S0-S3 and S6. Time evolution of the solute observations at the two verification wells #1, and #2 computed with the updated ensemble of source information parameters and  $\ln K$  after the assimilation of the solute observations of the first 60 time steps. The red line is the evolution of the concentration in the reference. The black lines correspond to the 5 and 95 percentiles of all realizations, and the green line corresponds to the median. The vertical dashed lines mark the end of the assimilation period.

iterations. The AAB( $\ln K$ ) and ESp( $\ln K$ ) are the smallest for S0 (r-EnKF), and the piezometric head and concentration predictions are also the best for S0. Only the ES-MDA with ten assimilation steps (S6) gives comparable results, although at a CPU cost 2.6 times larger than the r-EnKF.

However, when we analyze the reproduction of the contaminant source parameters, we 334 have already discussed Figure 5 showing that the ES-MDA is superior to the r-EnKF. This 335 observation is complemented by the results shown in Figure 11, in which boxplots of the 336 initial ensemble and the updated ensemble of the source parameters for the six scenarios are 337 shown. Some observations that can be derived from this figure are: the r-EnKF (S0) produces 338 good estimates for X, Y and T with a considerable reduction of uncertainty with respect 339 to the initial ensemble, while the estimates for  $\Delta T$  and M are somehow biased without a 340 large reduction of uncertainty; the ES (S1) is not effective, the spreads of the ensemble is 341 almost the same as for the initial ensemble prior to assimilation for all parameters; the ES-342 MDA starts to work well after four iterations, and gives the best results for ten iterations, 343 outperforming the r-EnKF, particularly for parameters X,  $\Delta T$  and M. The difficulty on 344 estimating  $\Delta T$  and M is due to the fact that several combinations of these two parameters 345 can result in very similar sets of observations, making more difficult their identification with 346 a reduction of uncertainty. The only way to solve this indetermination is the collection of 347 additional observations. This is precisely the reason why the ES-MDA with ten iterations 348 works better than the r-EnKF for these two parameters: the r-EnKF uses all observational 349 data only once in a piecewise way, whereas the ES-MDA uses all observation data altogether 350 ten times. 351

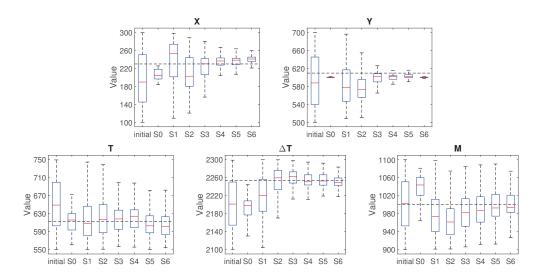


Figure 11: Scenarios S0-S6. Boxplots of the source location (X and Y), initial release time (T), release duration  $(\Delta T)$ , and mass-loading rate (M) computed with the initial parameters and with the updated parameters after 60 time steps. The dashed horizontal black line corresponds to the reference value.

#### 352 5. Summary and Discussion

The purpose of this paper is to analyze the ability of the ES-MDA for the identification of contaminant source parameters together with a spatially heterogeneous hydraulic conductivity field in comparison with the r-EnKF. The results show that the ES-MDA has the ability to estimate hydraulic conductivity field and identify the contaminant source parameters including source location, initial release time, release duration and mass-loading rate— with a proper number of iterations, besides, the results also indicate that these estimate parameters are good enough to provide good forecasts of solute concentrations and piezometric heads.

It is also worth pointing out that this is the first time that the ES-MDA is applied for contaminant source identification.

Furthermore, the comparison over all scenarios (including also the scenarios in the appendix) between the r-EnKF and the ES-MDA, shows that the ES-MDA performs better than the r-EnKF, especially for the identification of contaminant source parameters when using enough number of iterations. For the specific test done here, the ES-MDA starts to <sup>366</sup> outperform the r-EnKF after four iterations, needing almost the same computer time as that <sup>367</sup> for r-EnKF. The ES-MDA can perform even better using more iterations (and at a higher <sup>368</sup> computational cost). Part of the reason of the better performance of ES-MDA than of r-<sup>369</sup> EnKF is the fact that the number of observations is much larger for the ES-MDA, which is <sup>370</sup> specially important for the proper identification of mass-loading and release duration. These <sup>371</sup> two parameters are identified with large uncertainty by the r-EnKF.

It hovers over the whole paper whether an analysis on a single synthetic test case on 372 seven scenarios is sufficient to draw general conclusions about the comparison between the 373 r-EnKF and the ES-MDA. The answer is no, but drawing general conclusions was not the 374 purpose of this paper, its purpose was to test the newcomer ES-MDA against the r-EnKF in 375 a setting in which the r-EnKF had already proven to be quite effective. Given our extensive 376 experience with the application of the r-EnKF, we can forecast that a sensitivity study to 377 the number of observations will show that there is a threshold number below which it will be 378 impossible to identify the source; or that reducing the number of members of the ensemble 379 will require the use of localization and covariance inflation techniques to reach similar results, 380 with a threshold number of realizations below which identification will be impossible; or 381 that including a more uncertain prior distribution for the parameters describing the source 382 will have little impact to effectively identify the source beyond increasing the number of 383 assimilation steps. 384

While the just-mentioned sensitivity analyses are worth to carry out in a further study, there is an even more interesting issue that has not been addressed neither with the r-EnKF nor the ES-MDA, which is the analysis of more complex contamination events, such as non-punctual or multiple source ones. Addressing these events would require a thoughtful parameterization of the source.

#### <sup>390</sup> Appendix A. Results of scenarios S4 and S5

Results for scenarios S4 and S5 are displayed in Figures A.12 to A.15. The details are as follows: Figure A.12 shows the ensemble mean and ensemble variance of the updated  $\ln K$ ; Figure A.13 and A.14 show the 300th realization, ensemble mean and ensemble variance of piezometric heads and of the contaminant plume at the end of the 60th time step, respectively; Figure A.15 and A.16 show the time evolution of the piezometric heads and of solute concentrations at the two verification wells #1, and #2 computed with the updated source parameters and hydraulic conductivities.

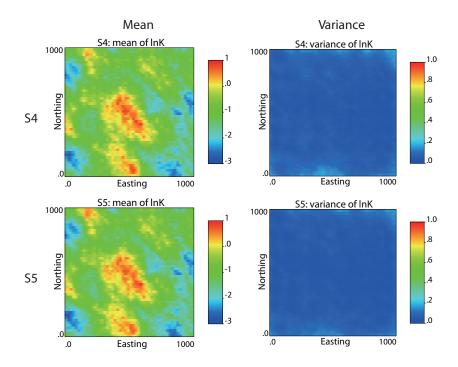


Figure A.12: Scenarios S3-S4. Ensemble mean (left column) and ensemble variance (right column) of updated log-conductivity realizations. (This figure complements Figure 4.)

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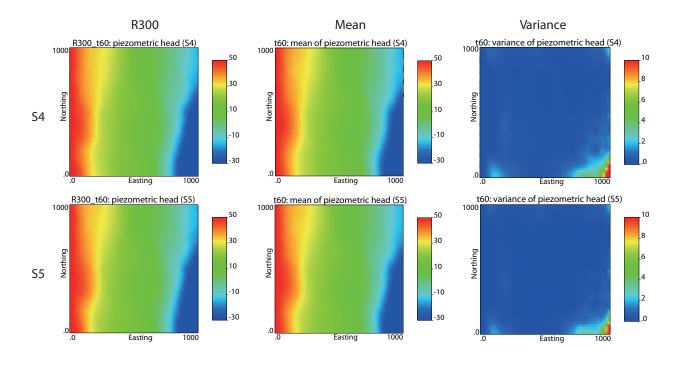


Figure A.13: Scenarios S4-S5. Piezometric heads computed with the updated parameters at the end of the 60th time step. From left to right, heads in realization #300; ensemble mean, and ensemble variance. (This figure complements Figure 6.)

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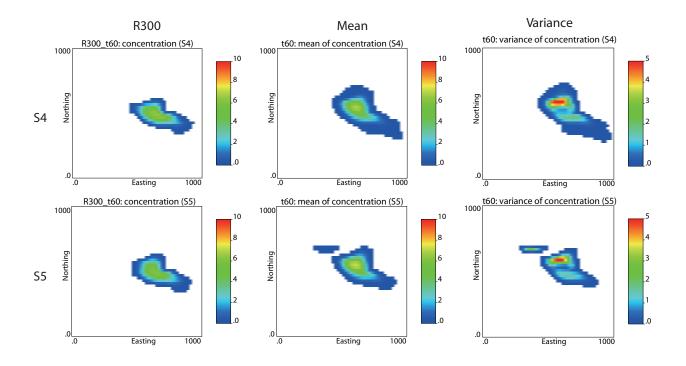


Figure A.14: Scenarios S4-S5. Contaminant plume computed with the updated parameters at the end of the 60th time step. From left to right, Contaminant plume in realization #300; ensemble mean, and ensemble variance. (This figure complements Figure 7.)

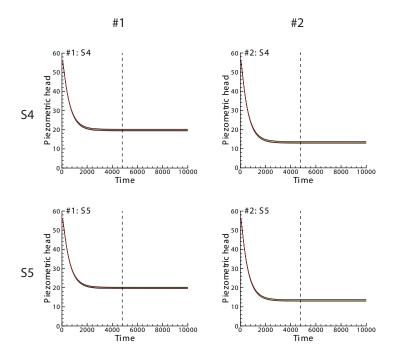


Figure A.15: Scenarios S4-S5. Time evolution of the piezometric heads at the two verification wells #1, and #2 computed with the updated ensemble of source information parameters at the end of the 60th time step. The red line is the evolution of the piezometric head in the reference. The black lines correspond to the 5 and 95 percentiles of all realizations, and the green line corresponds to the median. The vertical dashed lines mark the end of the assimilation period. (This figure complements Figure 9.)

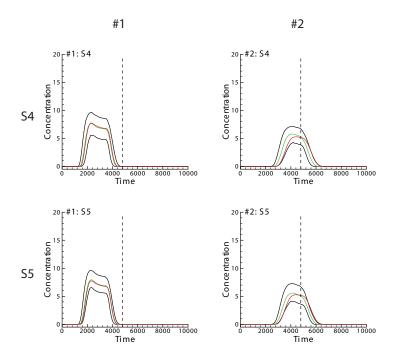


Figure A.16: Scenarios S4-S5. Time evolution of the solute concentrations at the two verification wells #1, and #2 computed with the updated ensemble of source information parameters at the end of the 60th time step. The red line is the evolution of the solute concentration in the reference. The black lines correspond to the 5 and 95 percentiles of all realizations, and the green line corresponds to the median. The vertical dashed lines mark the end of the assimilation period. (This figure complements Figure 10.)

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