# Joint identification of contaminant source and aquifer geometry in a sandbox experiment with the restart Ensemble Kalman filter

Zi Chen<sup>a</sup>, J. Jaime Gómez-Hernández<sup>a</sup>, Teng Xu<sup>a,\*</sup>, Andrea Zanini<sup>b</sup>

<sup>a</sup>Institute of Water and Environmental Engineering, Universitat Politècnica de València, Valencia, Spain <sup>b</sup>Dipartimento di Ingegneria e Architettura, Università degli Studi di Parma, Parma, Italy

# Abstract

Contaminant source identification is a key problem in handling groundwater pollution events. The ensemble Kalman filter (EnKF) is used for the spatiotemporal identification of a point contaminant source in a sandbox experiment, together with the identification of the position and length of a vertical plate inserted in the sandbox that modifies the geometry of the system. For the identification of the different parameters, observations in time of solute concentration are used, but not of piezometric head data since they were not available. A restart version of the EnKF is utilized because it is necessary to restart the forecast from time zero after each parameter update. The results show that the restart EnKF is capable of identifying both contaminant source information and aquifer-geometry-related parameters together with an uncertainty estimate of such identification.

*Keywords:* Inverse modeling; Observation error; Groundwater laboratory experiment; Stochastic hydrogeology

# <sup>1</sup> 1. Introduction

The problem of identifying a contaminant source in an aquifer using solute concentration data has been the subject of attention for many years (e.g., Atmadja and Bagtzoglou,

<sup>\*</sup>Corresponding author Email address: tenxu@posgrado.upv.es (Teng Xu )

2001; Michalak and Kitanidis, 2004; Bagtzoglou and Atmadja, 2005; Sun et al., 2006, and 4 references therein). Briefly, the proposed methods could be grouped into two categories: 5 optimization approaches and probabilistic approaches. The main difference between the two 6 approaches is that the optimization approaches cast the problem as a deterministic one in 7 which parameters are found that minimize a given objective function, whereas the probabilis-8 tic approaches cast the problem in a stochastic framework and the parameters to estimate 9 become random variables. In the first category, Gorelick et al. (1983) identified the ground-10 water pollution source information through an optimization model using linear programming 11 and multiple regression; Wagner (1992) employed a non-liner maximum likelihood method 12 to estimate source location and flux; Mahar and Datta (2000) used a nonlinear optimization 13 model for estimating the magnitude, location and duration of groundwater pollution sources 14 with binding equality constraints; Yeh et al. (2007) developed a hybrid approach, which 15 combines simulated annealing, tabu search and a three-dimensional groundwater flow and 16 solute transport model to solve the source identification problem; and Ayyaz (2010) utilized 17 a harmony search-based simulation-optimization model to determine the source location and 18 release histories by using an implicit solution procedure. In the second category, Bagtzoglou 19 et al. (1992) applied a particle method to estimate, probabilistically, source location and 20 spill-time history; Woodbury and Ulrych (1996) used a minimum relative entropy approach 21 to recover the release and evolution histories of a groundwater contaminant plume in a one-22 dimensional system; Neupauer and Wilson (1999) employed a backward location model based 23 on adjoint state method (BPM-ASM) to identify a contaminant source; Butera et al. (2013) 24 utilized a simultaneous release function and source location identification (SRSI) method to 25 identify the release history and source location of an injection in a groundwater aquifer; and 26 Koch and Nowak (2016) derived and applied a Bayesian reverse-inverse methodology to infer 27 source zone architectures and aquifer parameters. 28

The ensemble Kalman filter (EnKF), which could be included in the group of probabilistic approaches mentioned above, has recently addressed the problem of contaminant source

identification. The EnKF introduced by Evensen (2003) has gained much popularity in recent 31 years for its efficiency in solving inverse problems in different fields such as oceanography, 32 meteorology and hydrology (Houtekamer and Mitchell, 2001; Li et al., 2012a; Xu et al., 33 2013b). The advantages of the EnKF can be summarized as follows (Chen and Zhang, 2006; 34 Zhou et al., 2011): computational efficiency when compared with other inverse approaches, 35 easy integration with different forecast models, ability to account for model and observation 36 errors, and easy uncertainty characterization since the final outcome is always an ensemble 37 of realizations. In hydrogeology, the EnKF has been mainly applied for the identification of 38 aquifer parameters such as hydraulic conductivity or porosity (Li et al., 2012b; Xu et al., 39 2013a; Zhou et al., 2014; Xu and Gómez-Hernández, 2015, 2016a). Recently, Xu and Gómez-40 Hernández (2016b) demonstrated the possibility to apply the EnKF for the identification of a 41 contaminant source in a deterministic synthetic aquifer, and later Xu and Gómez-Hernández 42 (2018) showed that the method can be also applied for the simultaneous identification of 43 hydraulic conductivities and the parameters defining a contaminant source also in a synthetic 44 aquifer. 45

All the works mentioned above were tested in synthetic cases. Only a few works can be 46 found in the literature for laboratory or field cases. Woodbury et al. (1998) extended the 47 minimum relative entropy (MRE) method to recover the release history of a contaminant and 48 applied it to reconstruct the release history of a 1.4-dioxane plume observed at the Gloucester 49 Landfill in Ontario, Canada. Michalak (2003); Michalak and Kitanidis (2004) employed a 50 Bayesian inverse formulation to estimate the contaminant history of trichloroethylene (TCE) 51 and perchloroethylene (PCE) in an aquifer at the Dover Air Force Base, Delaware, a site that 52 had already been analyzed by Liu and Ball (1999) in the same context of source identifica-53 tion. Cupola et al. (2015b,a) compared the source location identification (SRSI) method to 54 the backward probability model based on the adjoint state method (BPM-ASM) with data 55 taken from a sandbox experiment. Zanini and Woodbury (2016) also used data from a sand-56 box experiment to apply an empirical Bayesian method combined with Akaike's Bayesian 57

Information Criterion (ABIC) to deduce the release history of a groundwater contaminant. 58 The main objective of this paper is to assess the performance of the restart EnKF (r-59 EnKF) for the identification of contaminant source parameters and aquifer geometry with 60 data from a sandbox experiment. The source parameters of interest are the release location, 61 release starting and ending times, and contaminant load, and regarding the geometry the 62 method should try to retrieve the position and length of a plate that is inserted about the 63 center of the sandbox and induces a deflection of the flowlines towards the bottom of the 64 sandbox. The state information assimilated by the r-EnKF is limited to concentration data 65 at a few observation points, since no piezometric head data were available. 66

The paper is organized as follows, first, the state equations and the fundamentals of the 67 r-EnKF will be recalled, second, the sandbox characteristics are described together with the 68 numerical model used to reproduce its behavior, third, the r-EnKF is tested with data from 69 a synthetic experiment that mimics the sandbox experiment with the aim to verify if the r-70 EnKF is capable of identifying the kind of parameters sought, and four, the r-EnKF is applied 71 with observation values taken from the sandbox experiment, the problems encountered are 72 analyzed, alternative approaches are discussed and the final results presented. The paper 73 ends with a summary and conclusions on the main findings. 74

## 75 2. Methodology

## <sup>76</sup> 2.1. Groundwater Flow and Solute Transport Equation

The sandbox will be modeled as a two-dimensional system in the XZ plane, where an inert contaminant spreads due to advection and dispersion under a steady-state flow. The dimension of the sandbox in the y direction is small enough to assume that the state variables are constant along any line for any given (x, z) value. The governing equations are:

$$\frac{\partial}{\partial x} \left( K_x \frac{\partial h}{\partial x} \right) + \frac{\partial}{\partial z} \left( K_z \frac{\partial h}{\partial z} \right) + w = 0 \tag{1}$$

$$\frac{\partial \left(\theta C\right)}{\partial t} = \nabla \cdot \left(\theta D \cdot \nabla C\right) - \nabla \cdot \left(\theta v C\right) - q_s C_s \tag{2}$$

where  $K_x$  and  $K_z$  are the principal components of the hydraulic conductivity tensor in the x 81 and z spatial coordinates respectively  $[LT^{-1}]$  which are assumed aligned with the coordinate 82 system of reference in the entire domain; h is the hydraulic head [L]; w represents distributed 83 sources or sinks  $[T^{-1}]$ ; t is time [T];  $\theta$  represents the porosity of the medium; C is dissolved 84 concentration  $[ML^{-3}]$ ;  $\nabla$  is the divergence operator;  $\nabla$  is the gradient operator; D represents 85 the hydrodynamic dispersion coefficient tensor  $[L^2T^{-1}]$ ; v is the flow velocity vector  $[LT^{-1}]$ 86 derived from the solution of the flow model;  $q_s$  represents volumetric flow rate per unit 87 volume of aquifer associated with a fluid source or sink  $[T^{-1}]$  and  $C_s$  is the concentration of 88 the source or sink  $[ML^{-3}]$ . 89

The flow equation is solved using MODFLOW (McDonald and Harbaugh, 1988), and the transport equation is solved using MT3DS (Zheng and Wang, 1999).

## 92 2.2. The Ensemble Kalman Filter

The ensemble Kalman filter was first introduced by Evensen (2003) to circumvent the 93 difficulty of propagating covariances in time in the original and extended Kalman filter 94 formulations. The restart EnKF (r-EnKF) has proven its capacity for contaminant source 95 identification in synthetic cases (Xu and Gómez-Hernández, 2016b, 2018); now, we propose 96 to test the r-EnKF in a sandbox experiment. For this specific case, there will be eight 97 parameters to identify, six related to the contaminant source, and two related to aquifer 98 geometry. In the first group, they are the contaminant source location (Xs, Zs), the injection 99 concentration Ic, the injection rate Ir, plus the starting Ts and ending Te release times. In 100 the second group, the algorithm will try to identify the position along the x direction Xb101 and the total depth Zb of a vertical plate inserted about the center of the sandbox to deflect 102 the flowlines. The rest of the parameters defining the flow and transport conditions in the 103 sandbox are not subject to identification and are equal to their observed values as explained 104

in the description of the experiment in the next section. The r-EnKF is shortly described
 next.

In the ensemble Kalman filter with extended state vector, we deal with two types of 107 variables, the system parameters subject of identification, of which there could be observa-108 tions or not, and the state of the system, of which there will be observations. The state 109 is forecasted in time solving the corresponding state equations, with the latest parameter 110 update, up to the specific time steps when observations are collected; these observations are 111 assimilated by the filter and serve to update the parameters and the state of the system. In 112 the restart filter, state variables are not updated, only system parameters are, because the 113 system state forecast for the next observation time is restarted from time zero to make sure 114 that the forecasted system state is fully coherent with the state equations, and, in our case, 115 with the updated **contaminant source**. (In the original implementation of the filter, both 116 state and parameters are updated, and the state system is forecasted from the last updated 117 state values using the last updated parameters.) The r-EnKF is an iterative algorithm that 118 cycles forecast and data assimilation (with the corresponding parameter update) until all 119 observations have been accounted for. The implementation of the r-EnKF for the identi-120 fication of the eight parameters described above can be summarized as follows (Evensen, 121 2003; Xu and Gómez-Hernández, 2016b): 122

123 1. Generate an initial ensemble of parameter values. An ensemble of  $N_e$  realizations 124 of eight-tuples of the parameters to be identified is generated. Parameter values are 125 drawn, independently, from uniform distributions defined between first-guess minimum 126 and maximum values—there are no restrictions on these uniform distributions, their 127 range can be wider or narrower than the one used in this paper, and they do not have 128 to necessarily contain the "real" value, they are simply used to initialize the algorithm. 129 We build  $N_e$  vectors  $S_i$  with the eight parameters for each realization:

$$S_i = [Xs_i, Zs_i, Xb_i, Zb_i, Ic_i, Ir_i, Ts_i, Te_i]^T$$
(3)

where i is the realization index and the superscript T stands for transpose.

2. Repeat for each system state observation time. Forecast the state. For each ensemble 131 member, forecast the system state, that is, the concentrations in the aquifer, for the 132  $t^{th}$  observation time using the values of the parameters from the last update (or the 133 initial parameters for the first observation time). In the original implementation of 134 the EnKF, the system state at the  $t^{th}$  observation time is forecasted based on the 135 concentrations at the  $(t-1)^{th}$  observation time and using the last updated parameters; 136 however, it is virtually impossible to account for an update of the source location or 137 the injection time unless the state equation is solved from time zero, thus the need to 138 restart the simulation from time zero (Xu and Gómez-Hernández, 2016b). The forecast 139 of concentrations is given by 140

$$C_i^f(t) = \psi \left[ C_0, S_i^a(t-1) \right],$$
(4)

where the superscripts f and a refer to forecasted and updated values after assimilation, respectively;  $\psi$  represents the numerical model that forecast, in time, concentrations, on a grid with  $N_m$  nodes;  $C_i$  is an  $N_m \times 1$  column vector containing the forecasted concentrations at all the discretization nodes of the numerical model for realization i;  $S_i^a$  is the vector with the last updated parameters;  $C_0$  is the initial contaminant concentration of the domain, which is the same for all realizations. The forecast of the parameters is simply

$$S_i^f(t) = S_i^a(t-1). (5)$$

3. Parameters update. First compute the parameter covariance through the ensemble of
 forecasted realizations

$$P_S^f(t) = \frac{1}{N_e} \sum_{i=1}^{N_e} \left\{ \left[ S_i^f(t) - \overline{S_i^f(t)} \right] \left[ S_i^f(t) - \overline{S_i^f(t)} \right]^T \right\}$$
(6)

150 with

$$\overline{S_i^f(t)} = \frac{1}{N_e} \sum_{i=1}^{N_e} S_i^f(t),$$
(7)

where  $P_S^f$  is an 8×8 matrix of parameter covariances and  $\overline{S_i^f(t)}$  is an 8×1 column vector of parameter averages. Then, compute the parameter-concentration cross-covariances but only with the forecasted concentration values that fall at concentration observation locations for time t (for the sake of simplicity, we will assume that observations are taken coinciding with some of the numerical model nodes, if not, there will be a need to provide a linear averaging procedure to estimate concentrations at observation locations from model concentration forecasts)

$$P_{SC}^{f}(t) = \frac{1}{N_e} \sum_{i=1}^{N_e} \left\{ \left[ S_i^{f}(t) - \overline{S_i^{f}(t)} \right] \left[ C_i^{f}(t) - \overline{C_i^{f}(t)} \right]^T \right\}$$
(8)

158 with

$$\overline{C_i^f(t)} = \frac{1}{N_e} \sum_{i=1}^{N_e} C_i^f(t),$$
(9)

where  $P_{SC}^{f}$  is an  $8 \times N_{o}$  matrix of parameter-concentration cross-covariances, with  $N_{o}$ being the number of nodes of the numerical model at which observations are taken at time step t, and  $\overline{C_{i}^{f}(t)}$  is an  $N_{o} \times 1$  column vector of average concentrations. Next, compute the  $8 \times N_{o}$  Kalman gain matrix K(t) as

$$K(t) = P_S^f(t) [P_{SC}^f(t) + R(t)]^{-1}$$
(10)

where R(t) is an  $N_o \times N_o$  diagonal observation error covariance matrix (implying that there is no correlation between observation errors) and proceed to update the parameter values, realization by realization by

$$S_i^a(t) = S_i^f(t) + K(t) \left[ d_i(t) - C_{io}^f(t) \right],$$
(11)

where  $d_i(t)$  is an  $N_o \times 1$  vector of observed concentrations (including observation errors with covariance given by R(t)) and  $C_{io}^f(t)$  is an  $N_o \times 1$  vector of forecasted concentrations.



## 170 3. Experimental Case

## 171 3.1. Description of the experiment

A single point pollution experiment was performed in a sandbox using sodium fluorescein 172 as tracer. The sandbox is built in plexiglass and has external dimensions of 120 cm  $\times$  14 173 cm  $\times$  70 cm as sketched in Fig. 1. The internal volume of 96 cm  $\times$  10 cm  $\times$  70 cm is 174 filled with constant-diameter spherical glass beads. There are two reservoirs at the edges of 175 the box imposing constant water levels of 60.7 cm and 53.6 m upstream and downstream, 176 respectively. An injector was set up at the upstream part of the sandbox at the location 177 indicated by a red square in the figure, and a plastic plate was vertically inserted inside 178 the glass beads in the middle of the sandbox, whose position and length is also shown in 179 the figure. The experimental equipment was placed in a dark box and a digital camera was 180 used to capture, every 5 s, the fluorescein luminosity within the rectangular zone of 85 cm 181 by 44 cm marked with a ticked rectangle in Fig. 1. The pictures were then processed and 182 the fluorescein luminosity transformed into concentrations after a calibration procedure, as 183 described by Citarella et al. (2015). In this case, eight different fluorescein concentrations 184 (C = 0; 2.5; 5; 10; 20; 25; 30; 35 mg/l) were used to calibrate and generate the luminosity-185 concentration curves in each picture pixel. The total experiment time lasted 1965 s, the 186 injection started at time 120 s and finished at time 1000 s. During the experiment, the rate 187 and concentration of the injection were also recorded. 188

It is very important to note that there are no piezometric head observations. The design of the tank did not allow for those observations. Had there been piezometric head data, they could have been assimilated in the filter and, without doubt, would have helped in improving
the identification (as shown by Xu and Gómez-Hernández (2018)).

## 193 3.2. Numerical Model

Since the thickness of the sandbox along the y axis is relatively small, we can assume 194 that the variability of piezometric heads and concentration along this direction is negligible. 195 Therefore, a two-dimensional groundwater flow and transport model in the XZ plane is 196 built. The upstream and downstream vertical boundaries are set as constant prescribed 197 piezometric head values, and the bottom boundary is impermeable while the top boundary 198 is the phreatic surface. The model corresponds to the yellowish area in Fig. 1, where 199 the coordinates of the four model corners are given. The tank is filled with homogeneous 200 spherical glass beads with a conductivity of 0.58 cm/s and a porosity of 0.37. The vertical 201 plastic plate was inserted at a distance of 52 cm from the left boundary and its length is of 202 42 cm. It is modeled as an impermeable barrier, which will deflect the flowlines towards the 203 bottom of the sandbox. The sandbox is discretized into 96 columns, one row, and 70 layers; 204 the size of each cell is  $(\Delta x, \Delta y, \Delta z) = (1, 10, 1)$  cm. The total simulation time is 1800 s and 205 is discretized into 90 uniform time steps. Citarella et al. (2015) evaluated the longitudinal 206 and transverse dispersivities of the spherical beads, resulting in values of 0.16 cm and 0.048 207 cm, respectively. The flow and transport parameters are collected in Table 1. 208

The release happens at coordinates (18.5 cm, 30.5 cm), with a concentration of 20 mg/l and an injection rate of 0.95 cm<sup>3</sup>/s. To start the ensemble Kalman filter 800 8-tuples of the source and plate parameters are generated from uniform distributions (not centered at the true values). The true values of the parameters to identify and the suspect range of the uniform distributions used to generate the initial ensemble are collected in Table 2.

## 214 4. Application

The objective of this work is to demonstrate the capacity of the r-EnKF for the identifi-215 cation of contaminant source information, including contaminant source location (Xs, Zs), 216 injection information (Ic, Ir) and release time (Ts, Te) together with the position and length 217 of the vertical plate (Xb, Zb), using concentration observations collected in a laboratory ex-218 periment. As a prior test, we analyze a synthetic case, in which the concentration data are 219 generated by the numerical model of the sandbox, therefore removing any modeling error 220 since the forward model used to forecast by the r-EnKF will coincide with the model used to 221 generate the observations. In the next section, we will redo the analysis using the laboratory 222 data, we will analyze the problems found and propose some solutions. 223

## 224 4.1. Synthetic Sandbox Test

In this case, we design two scenarios (S1, S2) with different number of observation wells 225 to evaluate the performance of the r-EnKF: scenario S1 with 20 observation wells, and 226 scenario S2 with 24 observation wells containing 4 additional wells (#21, #22, #23, #24) 227 located at the four corners of the suspect release area (see Fig. 1). The rationale for the 228 second scenario is that, after analyzing the first scenario, we felt that additional 229 information about the plume evolution was needed, and thus we decided to 230 add four wells around the suspect release zone. Such an addition will, indeed, 231 improve the characterization. In both scenarios, model error is neglected and we assume 232 that observation errors are uncorrelated with mean zero, and standard deviation of 0.1 mg/l. 233 Figure 2 and 3 show the time evolution of the ensemble mean and the ensemble variance, 234 respectively, of the updated state parameters for the two scenarios. Figure 4 shows the 235 evolution in time of the boxplots computed from the 800 ensemble members. After time 236 step 60, the convergence rate of the means and variances of the parameters 237 are less than 1% and 5%, respectively, all the parameters get close to the final 238 estimation and become stable. Notice also the sudden drop of the variance at a 239

given time step for most of the parameters. This drop is related to the activation 240 of new observation wells as time progresses, what implies that the amount of 241 information assimilated by the filter does not vary continuously in time, but 242 rather it increases stepwise, with steps occurring when new wells observe, for 243 the first time, the arrival of the solute plume. We can distinguish between the 244 parameters that are perfectly identified by an ensemble mean equal to the true value, and 245 practically zero variance, and those that are approximated closely but which are not exact 246 and present some residual uncertainty. In the first group, there are the position parameters 247 for the plate, Xb and Zb, plus the vertical location of the release source Zs, independently 248 of whether 20 or 24 data are used during the assimilation steps; in the second group are 249 the remaining parameters, which become more precise (mean closer to the true value) and 250 less uncertain (smaller variability) for S2 than for S1. The horizontal source location Xs251 is less sensitive to the concentration data, and only when the four additional data points 252 in the corners of the suspect release location are added the algorithm is able to provide a 253 good estimate for this parameter; similar comment can be made about the beginning Ts254 and end Te times of the release. The injection concentration Ic and injection rate Ir are 255 well identified by their median values, with smallest uncertainty for S2. These results are 256 consistent with the sensitivity of concentrations at the observation locations to changes in 257 the parameter values: concentration distributions are most sensitive to the position of the 258 plate, which affects the flow field, and the vertical release location, which affects the main 259 trajectory of the contaminant plume, but are less sensitive to the other parameters, for which 260 variations within the identified uncertainty ranges induce concentration changes of the same 261 order of magnitude as the observation errors. Also notice that the horizontal coordinate 262 of the release and the starting and ending release times are correlated for the purpose of 263 identifying their values (a displacement of the horizontal coordinate of the release could 264 be compensated with a displacement of its starting time), what also explains their larger 265 uncertainties. 266

These results prove that the r-EnKF could work for the identification of a contaminant source and of some parameters defining the geometry of the aquifer. The next step is to test the algorithm under more realistic conditions using observations obtained from a laboratory experiment.

#### 271 4.2. Laboratory Sandbox Test

The sandbox experiment was carried out as described previously. Figure 5 shows a 272 picture of the fluorescein plume at the 48th time step (840 s since the beginning of the 273 release) already transformed into concentration values and the position of the observation 274 points. The deflection of the flowlines induced by the vertical plate is clearly seen. Notice 275 that only a few observation piezometers will actually detect the plume breakthrough. Before 276 testing the r-EnKF, we performed a simulation of the concentration evolution using the 277 known release parameters and compared the predictions with the observed data. Figure 278 6 shows a comparison between observed and numerically predicted concentrations at five 279 observation locations (wells #7, #9, #10, #13, #22) through which the plume passes. As 280 can be seen, the reproduction is very good for the closest well #22, and it deteriorates with 281 the distance from the source, but not dramatically, except for well #9. For this well, the 282 beginning and ending times of the breakthrough curve are the same for predictions and 283 observations, but the mismatch in concentrations indicates either some error in the model 284 parameters or faulty observations. The predicted breakthrough curve in the farthest well, 285 though, is quite close to the observed one. In the application that follows we will analyze 286 different observation error distributions in an attempt to identify the source parameters by 287 the r-EnKF. 288

We have run the r-EnKF with three different magnitudes of the observation error, which will be referred to as R1, R2, and R3. In all three cases, the error mean is zero and its standard deviation is 0.5 mg/l for R1, 1.0 mg/l for R2, and 3.0 mg/l for R3. We must notice that in previous experiments, Cupola et al. (2015a) report an observation  $_{293}$  error with a standard deviation around 1 mg/l.

The hydraulic conductivity value of the beads, which is considered homogeneous in each 294 realization, is considered uncertain and drawn from a Gaussian distribution with a mean of 295 0.58 cm/s and a standard deviation of 0.05 cm/s. We have decided to introduce some 296 uncertainty on the beads hydraulic conductivity as a surrogate to model error. 297 The choice of a Gaussian distribution centered at the calibrated conductivity 298 value was arbitrary, any other distribution could have been used. Considering 299 that the differences in the results between including or not such an uncertainty 300 are minimal (and not reported here), we believe that the choice of the specific 301 distribution has little effect in the final outcome. 302

Fig. 7 shows the boxplots of the updated parameters at different time steps for the three 303 scenarios R1, R2, and R3. The results are not as good as for the synthetic case, for which the 304 observed concentrations were generated with the same numerical model used for the forecast 305 step in the Kalman filter. The first thing to note is that for scenario R1, the use of a small 306 observation error makes the r-EnKF to seek for source parameter values that can be far 307 from the true ones in order to produce concentrations that are close to the observed values, 308 and, particularly, the injection concentration and injection rate do not seem to converge to 309 a stable value after 90 time steps. The other parameters do reach a stable median, not as 310 close to the true values as for the synthetic case but close enough except for the horizontal 311 position of the vertical plate. 312

When the observation error is increased (scenario R2), the two main findings are that the two injection parameters now seem to reach a stable estimate (albeit with large uncertainty) with a median close to the true value, and that all parameters have a wider uncertainty range. The median estimate of the initial and ending release times is also closer to the true ones than in R1. The horizontal position of the vertical plate continues to be underestimated, as well as the length of the plate.

When the observation error is increased even more (scenario R3) the main effect is that

the final estimates have wide uncertainty estimates, and for some of the parameters it seems as if the concentration observations do not bring any added value since the boxplot width remains unaltered through the assimilation steps. The estimates of the parameters by their median is comparable to the results in R2, but their uncertainty is larger.

The predicted concentrations at three observation wells that were not used during the 324 assimilation step computed using the initial 8-tuples of parameters, and using the 8-tuples 325 obtained at the end of the three scenarios are shown in Figure 8. The figure shows the 326 true concentrations in the sandbox as a dotted blue line, each one of the 800 predicted 327 concentration breakthrough curves computed with the 8-tuples of the ensemble, along with 328 their median, as a red line, and their 90% confidence interval, as dashed lines. It can be 329 observed that, prior to assimilation (top row), concentration predictions were very scattered. 330 and that after the assimilation (bottom three rows, one for each scenario) the breakthrough 331 curves change substantially (compare, for instance, the median curves). For scenario R1, the 332 scatter of prediction curves is the smallest but recall that these wells were not used during 333 the assimilation, the updated parameters were biased because the algorithm tried to fit the 334 observed concentrations too closely and as a result, at the control wells, the prediction of 335 the true curves by the ensemble median is also biased, up to the point that the true curves 336 are outside the 90% confidence interval. For scenarios R2 and R3 the median curves for the 337 three wells have a smaller bias than for R1, and the main difference between R2 and R3 is 338 the same as for parameter prediction, the uncertainty is the widest for R3. The true curve 339 is in both cases within the 90% confidence interval of the predictions. 340

At this point, it seems that an observation error with a standard deviation of 1 mg/l was the most consistent with our observations and model. As mentioned above, this conclusion fits the findings by (Cupola et al., 2015b). Yet, we were concerned with the big discrepancy between predictions and observations at well #9, so we decided to rerun scenario R2 without using the data from this well. The results for this scenario, called R2b, are shown in Figure 9. When comparing this figure to the middle two rows in Figure 7 we can notice that there is some overall improvement in the estimation of the true parameters —particularly for the position parameters— by the median values of the ensemble without a significant change on their uncertainty. This improvement reinforces our suspicion that there could have been some problems in the data collection at well #9.

We also considered that there could be a problem with the tightness of the vertical plate 351 after its insertion in the sandbox. The plate was supposed to represent a perfectly 352 impermeable barrier, and no evidence of the contrary was observed during the 353 experiments, yet the contact between the plate and sandbox walls could have 354 had some small gaps, making the plate slightly permeable. Therefore, we decided 355 to rerun scenario  $R^2$  but assuming that the plate is slightly permeable, more precisely, with 356 a conductivity of two orders of magnitude smaller than the beads, this value was chosen 357 arbitrarily low since nothing was actually observed in the laboratory. The results 358 for the new scenario, referred to as R2c are shown in Figure 10. (Note that well #9 was kept 359 in this scenario.) The main difference of this run is that the estimate of the size of the vertical 360 plate by the median of the ensemble jumps from 40.5 cm to 44.2 cm (true value is 42.0 cm) 361 indicating that possibly the plate conductivity used in this scenario was too large and, as 362 a consequence, the algorithm enlarges the plate to reproduce the observed concentrations. 363 This result, while does not serve to justify that the tightness of the plate explains 364 the numerical model misfit, shows the impact that such permeability would have 365 in the estimation of the remaining parameters defining the plate. 366

We can conclude that the r-EnKF can be applied to a more realistic case of a homogeneous aquifer in a sandbox for the identification of a contaminant source and some geometry parameters. A proper evaluation of the observation errors is paramount, since attempting to match too closely the data may result in biased estimates of the parameters.

## 371 5. Summary and Conclusion

The main purpose of this paper was to test whether the restart ensemble Kalman filter, which had been successfully applied in synthetic experiments, could be applied to a more realistic case based on a sandbox experiment. The test focuses on the identification of the parameters defining a finite-pulse point injection of a solute, together with the position of a vertical plate that modifies the initial rectangular geometry of the sandbox.

As a preliminary step, we tested the r-EnKF in a synthetic case mimicking the sandbox. Under these very controlled conditions, the algorithm performs well, as expected. The main difference with previous synthetic analyses is that no piezometric head data were used during the assimilation step of the filter.

Then, the r-EnKF is tested using the data coming from the laboratory experiment. In this 381 case, the observations were not generated by a computer code nor we knew the observation 382 error magnitude. The analysis of the results show that using a too small observation error 383 variance results in more or less precise but biased estimates, both for the parameters subject 384 to identification and for the concentrations at control locations. When a larger observation 385 error (with a standard deviation of 1 mg/l) is introduced, estimates and predictions improve, 386 although with larger uncertainty. And finally, when the observation error is large, the results 387 worsen considerably. The removal of a suspicious observation well, the concentration of which 388 is always underestimated by our forecast model, improves the results, indicating that the 389 measurements from such well may need to be reconsidered. The changes observed after 390 making the vertical plate slightly permeable do not appear to justify the hypothesis that the 391 plate leaks. 392

The r-EnKF appears as a good algorithm for source identification in aquifers, yet it still needs further tests in closer-to-reality conditions. Currently, the sandbox has been replaced with a heterogeneous distribution of glass beads, and the challenge is to test the method in this new sandbox. Acknowledgements Financial support to carry out this work was received from the Spanish Ministry of Economy and Competitiveness through project CGL2014-59841-P, and from the Spanish Ministry of Education, Culture and Sports through a fellowship for the mobility of professors in foreign research and higher education institutions to the second author, reference PRX17/00150. The authors also would like to thank Università degli Studi di Parma for providing the experimental equipment and data.

## 403 **References**

- Atmadja, J., Bagtzoglou, A.C., 2001. State of the Art Report on Mathematical Methods for
  Groundwater Pollution Source Identification. Environmental Forensics 2, 205–214.
- Ayvaz, M.T., 2010. A linked simulation-optimization model for solving the unknown groundwater pollution source identification problems. Journal of Contaminant Hydrology 117,
  46–59.
- Bagtzoglou, A.C., Atmadja, J., 2005. Mathematical Methods for Hydrologic Inversion: The
  Case of Pollution Source Identification. Water Pollution 5, 65–96.
- <sup>411</sup> Bagtzoglou, A.C., Dougherty, D.E., Tompson, A.F.B., 1992. Application of particle methods
  <sup>412</sup> to reliable identification of groundwater pollution sources. Water Resources Management
  <sup>413</sup> 6, 15–23.
- <sup>414</sup> Butera, I., Tanda, M.G., Zanini, A., 2013. Simultaneous identification of the pollutant release
  <sup>415</sup> history and the source location in groundwater by means of a geostatistical approach.
  <sup>416</sup> Stochastic Environmental Research and Risk Assessment 27, 1269–1280.
- <sup>417</sup> Chen, Y., Zhang, D., 2006. Data assimilation for transient flow in geologic formations via
  <sup>418</sup> ensemble Kalman filter. Advances in Water Resources 29, 1107–1122.

- Citarella, D., Cupola, F., Tanda, M.G., Zanini, A., 2015. Evaluation of dispersivity coefficients by means of a laboratory image analysis. Journal of Contaminant Hydrology 172, 10–23.
- Cupola, F., Tanda, M.G., Zanini, A., 2015a. Contaminant release history identification in
  2-d heterogeneous aquifers through a minimum relative entropy approach. SpringerPlus
  4, 656.
- <sup>425</sup> Cupola, F., Tanda, M.G., Zanini, A., 2015b. Laboratory sandbox validation of pollutant
  <sup>426</sup> source location methods. Stochastic Environmental Research and Risk Assessment 29,
  <sup>427</sup> 169–182.
- Evensen, G., 2003. The Ensemble Kalman Filter: Theoretical formulation and practical
  implementation. Ocean Dynamics 53, 343–367.
- Gorelick, S.M., Evans, B., Remson, I., 1983. Identifying sources of groundwater pollution:
  An optimization approach. Water Resources Research 19, 779–790.
- Houtekamer, P.L., Mitchell, H.L., 2001. A Sequential Ensemble Kalman Filter for Atmospheric Data Assimilation. 0203058.
- Koch, J., Nowak, W., 2016. Identification of contaminant source architectures A statistical inversion that emulates multiphase physics in a computationally practicable manner.
  Water Resources Research 52, 1009–1025. 2014WR016527.
- Li, L., Zhou, H., Gómez-Hernández, J.J., Hendricks Franssen, H.J., 2012a. Jointly mapping hydraulic conductivity and porosity by assimilating concentration data via ensemble
  Kalman filter. Journal of Hydrology 428-429, 152–169.
- Li, L., Zhou, H., Hendricks Franssen, H.J., Gómez-Hernández, J.J., 2012b. Groundwater flow
  inverse modeling in non-MultiGaussian media: Performance assessment of the normalscore Ensemble Kalman Filter. Hydrology and Earth System Sciences 16, 573–590.

- Liu, C., Ball, W.P., 1999. Application of inverse methods to contaminant source identification
  from aquitard diffusion profiles at dover afb, delaware. Water Resources Research 35,
  1975–1985.
- Mahar, P.S., Datta, B., 2000. Identification of Pollution Sources in Transient Groundwater
  Systems. Water Resources Management 14, 209–227.
- McDonald, J.M., Harbaugh, A.W., 1988. A modular three-dimensional finite-difference flow
  model. Techniques of Water Resources Investigations of the U.S. Geological Survey, Book
  6, 586.
- <sup>451</sup> Michalak, A.M., 2003. A method for enforcing parameter nonnegativity in Bayesian in<sup>452</sup> verse problems with an application to contaminant source identification. Water Resources
  <sup>453</sup> Research 39, 1–14.
- <sup>454</sup> Michalak, A.M., Kitanidis, P.K., 2004. Estimation of historical groundwater contaminant
  <sup>455</sup> distribution using the adjoint state method applied to geostatistical inverse modeling.
  <sup>456</sup> Water Resources Research 40.
- <sup>457</sup> Neupauer, R.M., Wilson, J.L., 1999. Adjoint method for obtaining backward-in-time location
   <sup>458</sup> and travel time probabilities of a conservative groundwater contaminant. Water Resources
   <sup>459</sup> Research 35, 3389–3398.
- Sun, A.Y., Painter, S.L., Wittmeyer, G.W., 2006. A constrained robust least squares approach for contaminant release history identification. Water Resources Research 42, 1–13.
- Wagner, B.J., 1992. Simultaneous parameter estimation and contaminant source characterization for coupled groundwater flow and contaminant transport modelling. Journal of
  Hydrology 135, 275–303.
- <sup>465</sup> Woodbury, A., Sudicky, E., Ulrych, T.J., Ludwig, R., 1998. Three-dimensional plume source

reconstruction using minimum relative entropy inversion. Journal of Contaminant Hydrology 32, 131–158.

Woodbury, A.D., Ulrych, T.J., 1996. Minimum relative entropy inversion: Theory and application to recovering the release history of a groundwater contaminant. Water Resources
Research 32, 2671–2681.

<sup>471</sup> Xu, T., Gómez-Hernández, J.J., 2015. Inverse sequential simulation: A new approach for the
<sup>472</sup> characterization of hydraulic conductivities demonstrated on a non-Gaussian field. Water
<sup>473</sup> Resources Research 51, 2227–2242. 2014WR016527.

<sup>474</sup> Xu, T., Gómez-Hernández, J.J., 2016a. Characterization of non-Gaussian conductivities and
<sup>475</sup> porosities with hydraulic heads, solute concentrations, and water temperatures. Water
<sup>476</sup> Resources Research 52, 6111–6136.

<sup>477</sup> Xu, T., Gómez-Hernández, J.J., 2016b. Joint identification of contaminant source location,
<sup>478</sup> initial release time, and initial solute concentration in an aquifer via ensemble Kalman
<sup>479</sup> filtering. Water Resources Research .

<sup>480</sup> Xu, T., Gómez-Hernández, J.J., 2018. Simultaneous identification of a contaminant source
<sup>481</sup> and hydraulic conductivity via the restart normal-score ensemble Kalman filter. Advances
<sup>482</sup> in Water Resources 112, 106–123.

Xu, T., Gómez-Hernández, J.J., Zhou, H., Li, L., 2013a. The power of transient piezometric
head data in inverse modeling: An application of the localized normal-score EnKF with
covariance inflation in a heterogenous bimodal hydraulic conductivity field. Advances in
Water Resources 54, 100–118.

<sup>487</sup> Xu, T., Jaime Gómez-Hernández, J., Li, L., Zhou, H., 2013b. Parallelized ensemble Kalman
<sup>488</sup> filter for hydraulic conductivity characterization. Computers and Geosciences 52, 42–49.

- Yeh, H.D., Chang, T.H., Lin, Y.C., 2007. Groundwater contaminant source identification
  by a hybrid heuristic approach. Water Resources Research 43, 1–16.
- <sup>491</sup> Zanini, A., Woodbury, A.D., 2016. Contaminant source reconstruction by empirical Bayes
  <sup>492</sup> and Akaike's Bayesian Information Criterion. Journal of Contaminant Hydrology 185-186,
  <sup>493</sup> 74–86.
- <sup>494</sup> Zheng, C., Wang, P.P., 1999. MT3DMS: A Modular Three-Dimensional Multispecies Trans <sup>495</sup> port Model , 219.
- <sup>496</sup> Zhou, H., Gómez-Hernández, J.J., Hendricks Franssen, H.J., Li, L., 2011. An approach to
  <sup>497</sup> handling non-Gaussianity of parameters and state variables in ensemble Kalman filtering.
  <sup>498</sup> Advances in Water Resources 34, 844–864.
- <sup>499</sup> Zhou, H., Gómez-Hernández, J.J., Li, L., 2014. Inverse methods in hydrogeology: Evolution
  <sup>500</sup> and recent trends. Advances in Water Resources 63, 22–37.

Table 1: Parameters of the groundwater flow and transport model

Hydr. conduct., $K$	0.58  cm/s
Porosity, $\phi$	0.37
Long. disp., $\alpha_L$	$0.16 \mathrm{~cm}$
Transv. disp., $\alpha_T$	$0.048~\mathrm{cm}$

Table 2: Source and geometry parameters. True values and suspect ranges for the generation of the initial ensemble of realizations

\_

Parameter	Actual Value	Suspect Range
Xs (cm) - x-coordinate of source Zs (cm) - z-coordinate of source	$18.5 \\ 30.5$	16 - 25 23 - 32
Xb (cm) - x-coordinate of plate	50.5 52.5	50 - 59
Zb (cm) - plate length Ir (cm <sup>3</sup> /s) - injection rate	$\begin{array}{c} 42.5 \\ 0.95 \end{array}$	35 - 43 0.6 - 1.1
Ic (mg/l) - injection load	20	5-24
Ts (s) - starting release time $Te$ (s) - ending release time	$\begin{array}{c} 120 \\ 1000 \end{array}$	80 - 260 960 - 1140

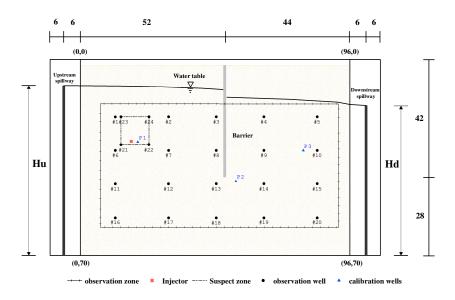


Figure 1: Sketch of the experimental device with indication of the upstream (Hu) and downstream (Hd) constant head boundaries. The ticked rectangle corresponds to the area captured by the camera in which concentrations will be monitored. Red dot is the release location. Dashed line around red dot indicates the release suspect location. Dimensions are in cm. Coordinates of the four corners of the flow and transport models are also shown.

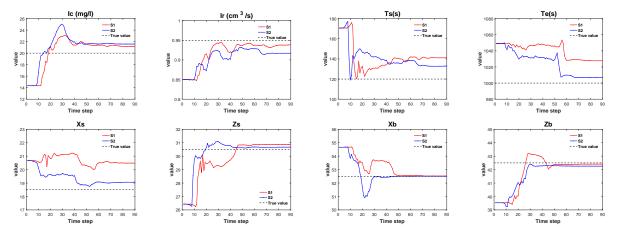


Figure 2: Time evolution of the ensemble mean of the 8 updated parameters, contaminant source location (Xs, Zs), plate position (Xb, Zb), injection information (Ic, Ir) and release time interval (Ts, Te) for scenarios S1 and S2

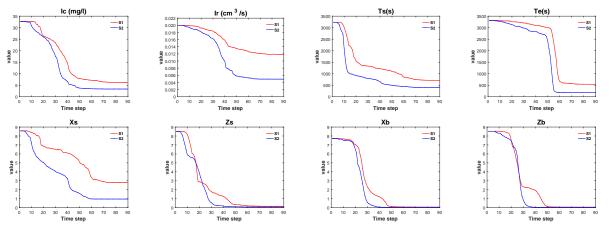


Figure 3: Time evolution of the ensemble variance for the same parameters and scenarios as in the previous figure.

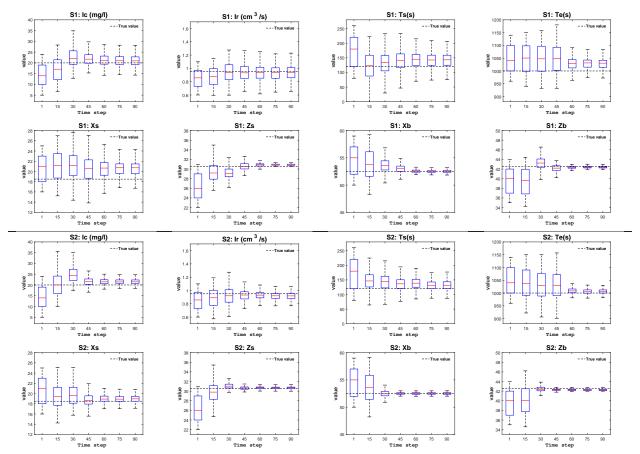


Figure 4: Boxplot of the 8 updated parameters at different time steps (1, 15, 30, 45, 60, 75, 90) for scenarios S1 and S2

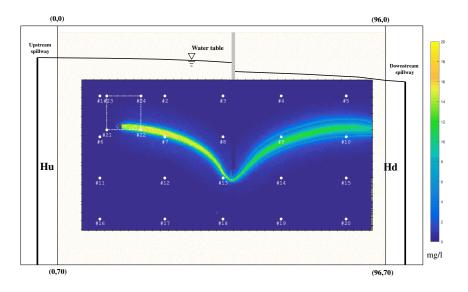


Figure 5: Fluorescein concentration field in the sandbox at the 48th time step. The area shown corresponds to the observation zone indicated in Fig. 1. The dash line shows the suspect zone for the injection and the white dots indicate the observation wells.

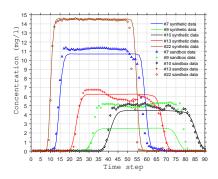


Figure 6: Fluorescein observed breakthrough curves at the observation wells located inside the plume and the curves computed from the numerical model



Figure 7: Boxplot of the 8 updated parameters at time steps 1, 15, 30, 45, 60, 75 and 90 for scenarios R1, R2 and R3

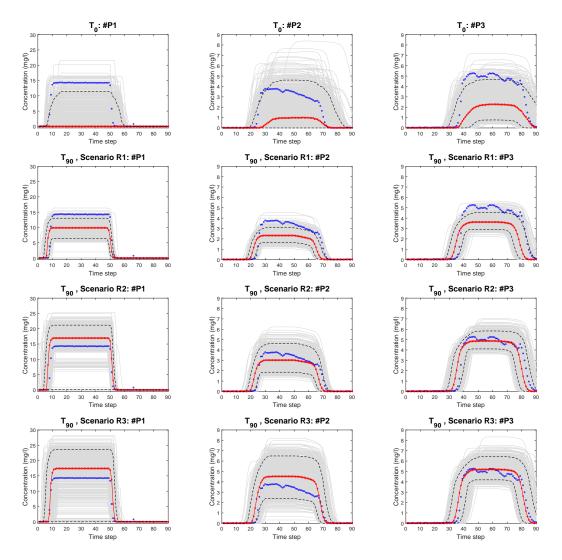


Figure 8: Breakthrough curves at control wells. The blue dots correspond to the curves in the sandbox experiment. The thin gray lines are the curves for all 800 realizations; they are summarized by their median (red diamond lines) and their 5 and 95 percentiles (black dash lines).

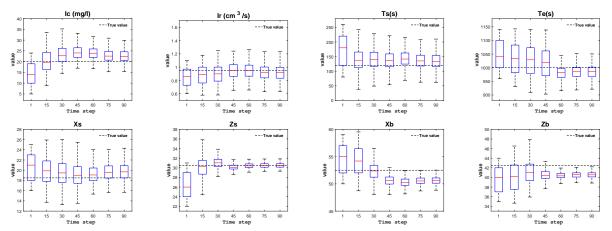


Figure 9: Boxplot of the 8 updated parameters in scenario R2b at different time steps (1, 15, 30, 45, 60, 75, 90)

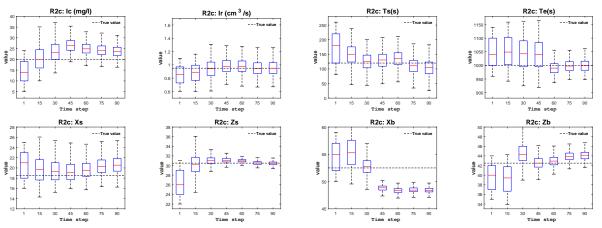


Figure 10: Boxplot of the 8 updated parameters in scenario R2c at different time steps (1, 15, 30, 45, 60, 75, 90)